

CO₂ INJECTION IN DEEP SALINE SLOPING AQUIFERS THROUGH A VERTICAL WELL

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INTRODUCTION

Anthropogenic CO₂ emissions are expected to continue to increase worldwide in the next decades. This could dramatically increase CO₂ concentrations in the atmosphere. One potential mitigation action is CO₂ permanent storage in deep saline aquifers. CO₂ is injected as a supercritical fluid (temperature and pressure higher than 31.1 °C and 7.38 MPa, respectively) so as to obtain a relatively high density that minimizes the volume occupied by this greenhouse gas. CO₂ density is highly dependent on temperature and pressure because of its high compressibility (Vilarrasa *et al.*, 2010a), adopting a wide range of values (450-800 kg/m³). Since CO₂ is lighter than brine, flow is affected by buoyancy.

Buoyancy produces an upslope migration of the CO₂ bubble in sloping aquifers. The post-injection fate of CO₂ in sloping aquifers has been investigated (e.g. Hesse *et al.*, 2008; Elenius *et al.*, 2010) but without considering the actual shape of the CO₂ bubble at the end of injection. However, the shape of the CO₂ bubble at the end of injection affects its posterior evolution (McMinn & Juanes, 2009). Gasda *et al.* (2008) studied the effect of a slope up to 1° during the injection period using a 2D model, which represents injection through a horizontal well. However, CO₂ injection through a vertical well, which implies a 3D geometry, has not yet been investigated.

CO₂ will be injected in the Hontomín CO₂ pilot storage site, Burgos (Spain). Hontomín is the site for the CO₂ storage Technology Demonstration Plant (TDP) of the Compostilla OXYCFB300 project, operated by Energy City Foundation (CIUDEN). CO₂ will be injected through a vertical well in a flank of a dome-like structure with a slope close to 20° at a depth around 1450 m. CO₂ injection tests are aimed to gain knowledge on trapping mechanisms and CO₂ bubble and pressure evolution. Pressure evolution is important for assessing the caprock integrity and to avoid the open up of preferential paths for CO₂ leakage (Vilarrasa *et al.*, 2010b, 2011). These processes are relatively well known in horizontal aquifers, but a high uncertainty still exists in sloping aquifers.

IMPLEMENTATION OF CO₂ MODULE IN CODE_BRIGHT

CODE_BRIGHT has been adapted to simulate CO₂ injection. To do so, an equation of state for CO₂ has been implemented. CO₂ density follows a Redlich-Kwong equation of state type with the parameters proposed by Spycher *et al.* (2003). CO₂ viscosity has been taken from Altunin & Sakhabetdinov (1972). These expressions for CO₂ density and viscosity are valid regardless of the temperature and pressure conditions (Figure 1).

Additionally, water density increase due to CO₂ dissolution has been implemented following Garcia (2003).

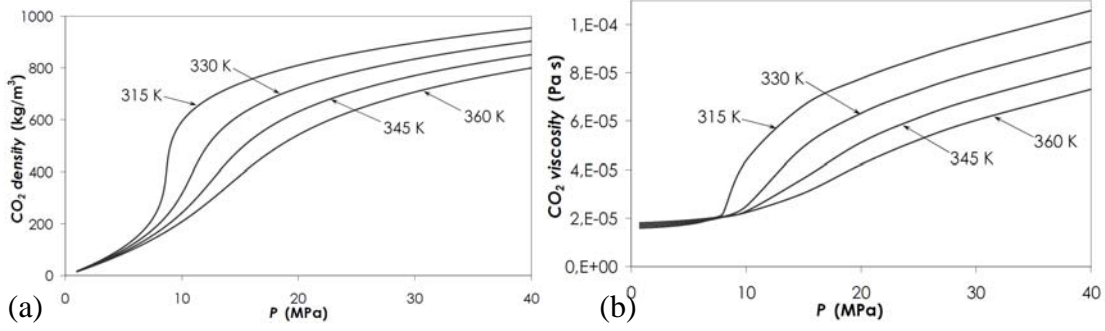


Figure 1. (a) CO₂ density and (b) viscosity as a function of pressure for several temperatures.

MODEL SETUP

To model CO₂ injection in a sloping homogeneous aquifer, we use a 3D model. Since we are interested in simulating CO₂ injection tests of 100 t of CO₂ during 1 week, we model half of the domain, 200x70x10 m³, by making use of the symmetry in the direction of maximum slope. The upslope and downslope boundaries are treated as constant head boundaries. The aquifer permeability is 10⁻¹³ m² and its porosity 0.1. The slope of the aquifer ranges from 0° to 20°. The mesh is structured, made of hexahedrons of 1 m in side. We use the finite element numerical code CODE_BRIGHT (Olivella *et al.*, 1994, 1996) modified for CO₂ injection.

GRAVITY NUMBER

Consider the injection of compressible CO₂ in a deep homogeneous horizontal confined brine aquifer through a vertical well. Mass conservation of these two fluids can be expressed as (Bear, 1972)

$$\frac{\partial(\phi S_{\alpha} \rho_{\alpha})}{\partial t} + \nabla \cdot (\rho_{\alpha} \mathbf{q}_{\alpha}) = 0, \quad (1)$$

where ϕ is porosity, t is time, S_{α} is the saturation of the α -phase, ρ_{α} its density, \mathbf{q}_{α} its volumetric flux and α is either w brine or c CO₂. Momentum conservation is expressed using Darcy's law

$$\mathbf{q}_{\alpha} = -\frac{kk_{\alpha}}{\mu_{\alpha}} (\nabla p_{\alpha} + \rho_{\alpha} g \nabla z), \quad (2)$$

where k is intrinsic permeability, k_{α} is relative permeability of the α -phase, μ_{α} its viscosity, p_{α} its pressure, g is gravity and z the vertical coordinate.

To quantify the relative influence of buoyancy we define a gravity number, N , as the ratio of gravity to viscous forces. We use the definition of Vilarrasa *et al.* (2010a) but adopted for sloping aquifers

$$N = \frac{2\pi r_{ch} b k \Delta \rho g \sin \theta \rho_{ch}}{\mu_c Q_m}, \quad (3)$$

where b is the aquifer thickness, $\Delta\rho$ is the difference between the fluids density, θ is the slope of the aquifer, Q_m is the mass flow rate and the subindex ch denotes a characteristic variable.

NUMERICAL SIMULATIONS

We model the injection of CO₂ in a sloping aquifer through a vertical well. This enables to simulate a realistic evolution of the CO₂ bubble, which starts to advance through the top of the aquifer at the beginning of injection due to buoyancy (Figure 2). The CO₂ bubble advances both laterally and downwards as CO₂ pressure increases. The thickness of the aquifer that gets desaturated depends mainly on the aquifer transmissivity and CO₂ flow rate. Once CO₂ injection stops, the CO₂ bubble is displaced upwards due to buoyancy (Figure 2). This produces the migration of the CO₂ bubble through the top of the aquifer. This might be problematic if solubility and capillary trapping do not completely trap the CO₂ in free phase before it reaches a spillpoint (Gasda *et al.*, 2008; Elenius *et al.*, 2010). The length that the CO₂ will reach depends on the gravity number.

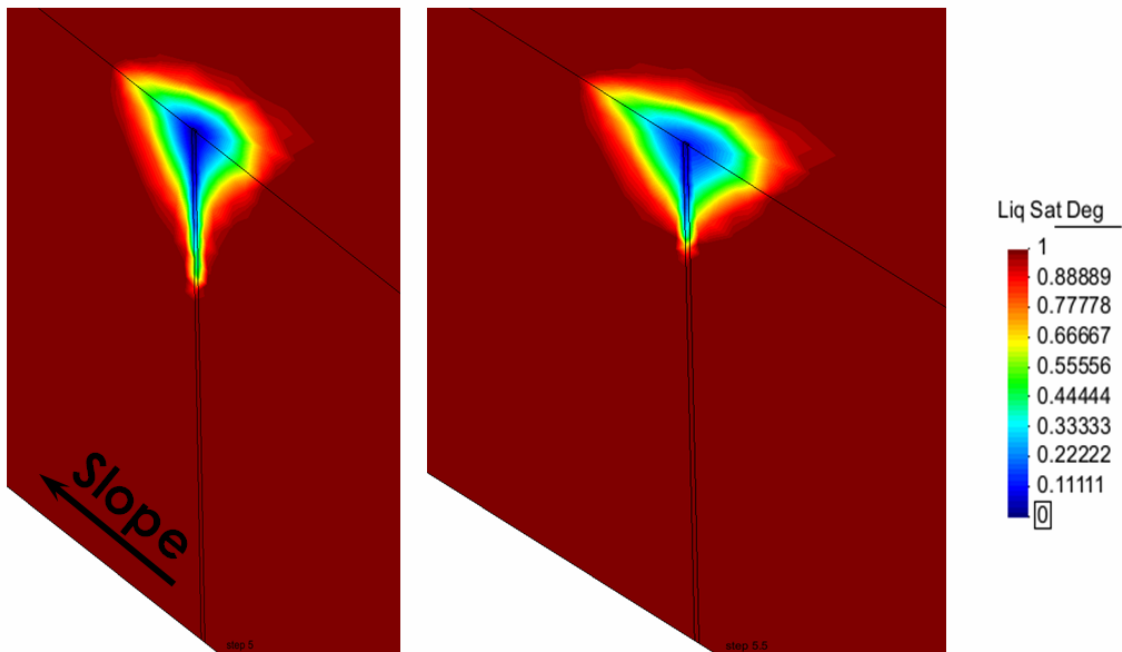


Figure 2. CO₂ bubble after 4 days of a 1 kg/s injection (left) and 0.5 days after injection stops (right). Note that the CO₂ bubble does not occupy the whole thickness of the aquifer (100 m thick). Note also that the CO₂ bubble is displaced upwards and dispersed laterally due to buoyancy once injection stops, advancing preferentially upslope.

CONCLUSIONS

CODE_BRIGHT has been successfully modified to simulate CO₂ injection in deep saline aquifers. The implemented density and viscosity functions are general and cover the whole range of depths and geothermal gradients of potential aquifers. CO₂ injection in sloping aquifers has been simulated using a 3D geometry.

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